

4-28-2016

Contributions of organic and inorganic matter to sediment volume and accretion in tidal wetlands at steady state

James T. Morris
University of South Carolina

Donald C. Barber
Bryn Mawr College

John C. Callaway
University of San Francisco

Randy Chambers
College of William and Mary

Scott C. Hagen
Louisiana State University

See next page for additional authors

Follow this and additional works at: https://digitalcommons.fiu.edu/fce_lter_journal_articles

 Part of the [Life Sciences Commons](#)

Recommended Citation

Morris, J. T. et al. (2016), Contributions of organic and inorganic matter to sediment volume and accretion in tidal wetlands at steady state, *Earth's Future*, 4, 110–121, doi:10.1002/2015EF000334.

This material is based upon work supported by the National Science Foundation through the Florida Coastal Everglades Long-Term Ecological Research program under Cooperative Agreements #DBI-0620409 and #DEB-9910514. Any opinions, findings, conclusions, or recommendations expressed in the material are those of the author(s) and do not necessarily reflect the views of the National Science Foundation.

This work is brought to you for free and open access by the FCE LTER at FIU Digital Commons. It has been accepted for inclusion in FCE LTER Journal Articles by an authorized administrator of FIU Digital Commons. For more information, please contact dcc@fiu.edu, jkrefft@fiu.edu.

Authors

James T. Morris, Donald C. Barber, John C. Callaway, Randy Chambers, Scott C. Hagen, Charles S. Hopkins, Beverly J. Johnson, Patrick Megonigal, Scott C. Neubauer, Tiffany Troxler-Gann, and Cathleen Wigand



RESEARCH ARTICLE

10.1002/2015EF000334

Contributions of organic and inorganic matter to sediment volume and accretion in tidal wetlands at steady state

Special Section:

Integrated field analysis & modeling of the coastal dynamics of sea level rise in the northern Gulf of Mexico

Key Points:

- Sediment bulk density can be largely explained by its organic matter concentration
- Bulk volume approximates the summed volumes of the mineral and organic fractions
- Upper bound of steady-state accretion defines sustainability of coastal marshes under SLR

Supporting Information:

- Supporting Information S1

Corresponding author:

J. T. Morris, morris@biol.sc.edu

Citation:

Morris, J. T. et al. (2016), Contributions of organic and inorganic matter to sediment volume and accretion in tidal wetlands at steady state, *Earth's Future*, 4, 110–121, doi:10.1002/2015EF000334.

Received 27 OCT 2015

Accepted 14 MAR 2016

Accepted article online 21 MAR 2016

Published online 28 APR 2016

James T. Morris¹, Donald C. Barber², John C. Callaway³, Randy Chambers⁴, Scott C. Hagen⁵, Charles S. Hopkinson⁶, Beverly J. Johnson⁷, Patrick Megonigal⁸, Scott C. Neubauer⁹, Tiffany Troxler¹⁰, and Cathleen Wigand¹¹

¹Department of Biological Sciences, Belle W. Baruch Institute for Marine & Coastal Sciences, University of South Carolina, Columbia, South Carolina, USA, ²Department of Geology, Bryn Mawr College, Bryn Mawr, Pennsylvania, USA, ³Department of Environmental Science, University of San Francisco, San Francisco, California, USA, ⁴Keck Environmental Field Laboratory, College of William and Mary, Williamsburg, Virginia, USA, ⁵Department of Civil & Environmental Engineering and Center for Computation & Technology, Louisiana State University, Baton Rouge, Louisiana, USA, ⁶Department of Marine Sciences, University of Georgia, Athens, Georgia, USA, ⁷Department of Geology, Bates College, Lewiston, Maine, USA, ⁸Smithsonian Environmental Research Center, Edgewater, Maryland, USA, ⁹Department of Biology, Virginia Commonwealth University, Richmond, Virginia, USA, ¹⁰Department of Biological Sciences, Southeast Environmental Research Center, Florida International University, Miami, Florida, USA, ¹¹Atlantic Ecology Division, U.S. EPA, Office of Research and Development, National Health and Environmental Effects Research Laboratory, Narragansett, Rhode Island, USA

Abstract A mixing model derived from first principles describes the bulk density (BD) of intertidal wetland sediments as a function of loss on ignition (LOI). The model assumes that the bulk volume of sediment equates to the sum of self-packing volumes of organic and mineral components or $BD = 1/[(LOI/k_1) + (1-LOI)/k_2]$, where k_1 and k_2 are the self-packing densities of the pure organic and inorganic components, respectively. The model explained 78% of the variability in total BD when fitted to 5075 measurements drawn from 33 wetlands distributed around the conterminous United States. The values of k_1 and k_2 were estimated to be $0.085 \pm 0.0007 \text{ g cm}^{-3}$ and $1.99 \pm 0.028 \text{ g cm}^{-3}$, respectively. Based on the fitted organic density (k_1) and constrained by primary production, the model suggests that the maximum steady state accretion arising from the sequestration of refractory organic matter is $\leq 0.3 \text{ cm yr}^{-1}$. Thus, tidal peatlands are unlikely to indefinitely survive a higher rate of sea-level rise in the absence of a significant source of mineral sediment. Application of k_2 to a mineral sediment load typical of East and eastern Gulf Coast estuaries gives a vertical accretion rate from inorganic sediment of 0.2 cm yr^{-1} . Total steady state accretion is the sum of the parts and therefore should not be greater than 0.5 cm yr^{-1} under the assumptions of the model. Accretion rates could deviate from this value depending on variation in plant productivity, root:shoot ratio, suspended sediment concentration, sediment-capture efficiency, and episodic events.

1. Introduction

Tidal wetlands have survived in place within the intertidal zone for millennia of rising sea level through biophysical processes that lead to the accumulation of mineral and organic matter, thereby increasing soil volume and surface elevation. Long-term wetland stability requires that the soil surface elevation must increase at a rate similar to the local rate of relative sea-level rise (SLR), i.e., they must be in equilibrium with the sea level. Many contemporary tidal wetlands formed during the late Holocene and accumulated up to several meters of deposits at a time when rates of SLR were relatively slow [Redfield and Rubin, 1962; Kelley et al., 1988; Engelhart and Horton, 2012; Engelhart et al., 2015]. With accelerating rates of relative SLR [Vermeer and Rahmstorf, 2009; Horton et al., 2014] and many examples of tidal wetlands converting to open water [Craft et al., 2009], there is an increasing need to understand the limits of the biophysical processes that expand soil volume and elevation.

Wetland elevation and its rate of change represent a balance between local rates of accretion and erosion as these factors ultimately determine the net change in sediment volume per unit area. Sediment volume decreases as a consequence of organic matter decomposition, compaction, and erosion. Sediment volume increases with additions of refractory root and rhizome tissue and deposition onto the sediment surface of

© 2016 The Authors.

This is an open access article under the terms of the Creative Commons Attribution-NonCommercial-NoDerivs License, which permits use and distribution in any medium, provided the original work is properly cited, the use is non-commercial and no modifications or adaptations are made.

mineral and refractory organic particles, which in turn is influenced by aboveground plant biomass. Empirical studies [e.g., *Nyman et al.*, 1993; *Turner et al.*, 2000; *Neubauer*, 2008] and simulation models [e.g., *Kirwan et al.*, 2010] have illustrated the interactive roles of organic matter and mineral inputs on wetland accretion rates.

The annual surface deposition of mineral and organic material creates a lamination that can be conceptualized as a sediment cohort [*Morris and Bowden*, 1986], the volume of which will change over time as a consequence of compaction and the subsequent ingrowth and turnover of roots and rhizomes. The final volume of a cohort is achieved when its organic matter content has stabilized, a process that occurs over many years of burial by generations of younger cohorts until the cohort has reached a depth below the root zone [*Morris and Bowden*, 1986; *Davis et al.*, 2015]. When the wetland is in equilibrium with a constant rate of SLR (i.e., the system is in steady state), the final mass of a cohort and its bulk density will determine the rate of vertical accretion. Many studies have described a variety of transient perturbations—including nutrient pollution, changing atmospheric chemistry, abrupt sea-level change, and watershed development—that can alter or disrupt steady-state accretion processes of tidal wetlands [*Langley et al.*, 2009; *Kirwan et al.*, 2011; *Deegan et al.*, 2012; *Kirwan and Megonigal*, 2013; *Mariotti and Fagherazzi*, 2013; *Weston*, 2014; *Ratliff et al.*, 2015]. Nevertheless, in light of the persistence of marsh habitat over centuries of rising sea level [*Kemp et al.*, 2011, 2013], wetland survival must be viewed in the context of the steady state. Transient perturbations return to equilibrium with mean sea level, or they lead to wetland loss and a state change.

A central requirement in many empirical and mechanistic studies of wetland accretion is the need to convert mineral and organic mass to a volume (and therefore thickness) of sediment. The analysis presented here was motivated by the simple objective of describing the physical relationship between sediment dry bulk density (BD) and organic matter concentration expressed as loss on ignition (LOI) in coastal wetlands. A number of publications have explored this relationship for upland or wetland soils, using either empirical models or first principles [*Jeffrey*, 1970; *Stewart et al.*, 1970; *Adams*, 1973; *Harrison and Bockock*, 1981; *Gosselink et al.*, 1984; *Grigal et al.*, 1989; *Honeysett and Ratkowsky*, 1989; *Manrique and Jones*, 1991; *Federer et al.*, 1993; *Turner et al.*, 2000; *Neubauer*, 2008; *Périeré and Ouime*, 2008]. Common limitations of this previous work include datasets that are limited geographically, by a narrow range of possible soil conditions, and by sample size. Herein, we extend these previous analyses to a wide sample of herbaceous and forested tidal wetland ecosystems (>5000 samples across 33 riverine and estuarine systems) from the coastal region of the conterminous United States. The present study tests a quantitative mixing model wherein the proportions and compositions of the two end-member components, organic and inorganic material, govern the density and volume of tidal wetland sediments and therefore influence steady-state rates of vertical accretion in tidal wetlands. The results suggest a simple and robust rule for setting upper limits on steady-state vertical accretion rates in coastal wetlands, a rule that is defensible from first principles.

2. Data and Methods

2.1. Sampling Protocol

We compiled a database of 5075 samples of sediment BD and LOI from 33 wetland sites in estuaries in Maine, Massachusetts, New York, Rhode Island, Delaware, Maryland, Virginia, North and South Carolina, Florida, Alabama, Mississippi, Louisiana, and California. The data were either unpublished and provided by the authors or published and publically available.

The methods for estimating bulk density and LOI in sediment vary in details, such as the operation, type, and dimension of the coring device, the volume of samples analyzed, etc. Despite this variation, they all share a common protocol of drying samples of known wet volume, weighing the dried samples, combusting the samples in a furnace, and reweighing to obtain the weight loss on ignition (LOI) [e.g., *Dean*, 1974; *Percival and Lindsay*, 1997]. There are numerous opportunities for error, including compression of the sediment when taking the core and/or when extracting the core from the core tube, imprecise sectioning of the core into known volumes, variation in drying temperature, furnace temperature, and presence of salts that precipitate when the pore water is evaporated from the sample [*Heiri et al.*, 2001]. We think that compression during the coring operation is the most significant source of error, but to our knowledge, the relative sizes of these errors have not been evaluated.

2.2. Statistics

The Marquardt-Levenberg method [Levenberg, 1944] of the SAS 9.4 PROC MODEL was used to estimate the unknown parameters in several alternative nonlinear models describing the relationship between sediment BD and LOI. PROC MODEL is an iterative procedure that finds the combination of parameters that minimizes the sum of squares of residuals. The procedure reports an adjusted R^2 value for the model being fitted and significance tests with associated probabilities that the parameters differ significantly from zero. Nonlinear regression analysis has an advantage over linear regression, which requires that nonlinear data must first be linearized. Linear transformation introduces a bias by changing the relative weights of individual observations. Additional tests were performed (PROC REG, PROC GLM) with log-transformed data to test for differences across sites, including homogeneity of slopes and LOI-by-site interactions. A massive dataset ($N > 6000$) from the Louisiana Coastwide Reference Monitoring System (CRMS) was randomly subsampled using PROC SURVEYSELECT to bring it in line with the number of samples from other sites and to prevent a deltaic bias.

3. Results

3.1. Site Characteristics

The 33 sites analyzed (Table 1) represented varied climatic conditions, tidal amplitudes, soil compositions, and species. Most samples are from the near surface (50 cm), except for cores from Sand Hill, NC that were nearly 3 m in length and cores from Maine, Patuxent River, MD, and Otter Point, MD, where 1–2 m of sediment was sampled. Sediments were predominately siliciclastic, with varying amounts of organic matter, but also included calcareous sediments from the coastal Everglades (FCE). The sites spanned 14 degrees of latitude. Across sites, mean sediment LOI ranged from 0.12 g g^{-1} at Grand Bay, MS; China Camp, CA; and Otter Creek, MD to 0.79 g g^{-1} at GCREW near Edgewater, MD. Mean dry bulk density ranged from 0.12 g cm^{-3} at GCREW to 0.78 g cm^{-3} at North Inlet, SC.

The dominant species vary among sites, with *Spartina patens* and *Spartina alterniflora* most common in northeastern and southeastern estuaries, respectively; *Spartina foliosa* and *Schoenoplectus* in the San Francisco estuary; a mix of *Spartina* species and brackish species (*S. cynosuroides* and/or *Juncus roemerianus*) in the Gulf Coast and North Carolina estuaries; and the red mangrove *Rhizophora mangle* in several Everglades sites. The tidal freshwater wetlands typically contained a diverse mixture of herbaceous species, including *Peltandra virginica*, *Pontederia cordata*, *Zizania aquatica*, and *Zizaniopsis miliacea*, with some locations also containing trees (e.g., *Nyssa aquatica*, *Taxodium distichum*). There was a weak, but insignificant ($p = 0.25$) decline of mean LOI by site with increasing latitude (-0.01 per degree) and a weak decline in LOI with tidal amplitude (-0.002 per cm, $R^2 = 0.27$, $p = 0.05$).

Tests of homogeneity of slopes of log-normalized data were significant ($p < 0.0001$). Likewise, an analysis of covariance established that LOI varies among sites: the independent variables site, $\ln(\text{LOI})$, and the $\ln(\text{LOI}) \times \text{site}$ interaction term were all significant ($p < 0.0001$) for type I and type III errors, meaning that there is no reasonable chance that these samples behave in the same way with respect to the dependence of log-transformed BD on log-transformed LOI.

3.2. Model 1 Power Function

The first model tested (Figure 1) is one that is commonly used to estimate bulk density (BD) from LOI measurements. It is an empirical power function of the form

$$\text{BD} = a\text{LOI}^b \quad (1)$$

where a and b are empirical constants. The equation was fitted with an R^2 of 0.73. Parameter $a = 0.141 \pm 0.002$ ($p < 0.0001$) and $b = -0.59 \pm 0.004$ ($p < 0.0001$).

3.3. Model 2 Power Function With Intercept

This model differs from the first only in having a nonzero intercept c . It is an empirical power function of the form

$$\text{BD} = a\text{LOI}^b + c \quad (2)$$

where a , b , and c are empirical constants. The equation was fitted with an R^2 of 0.77. Parameter $a = 0.569 \pm 0.03$ ($p < 0.0001$), $b = -0.31 \pm 0.01$ ($p < 0.0001$), and $c = -0.55 \pm 0.04$ ($p < 0.0001$).

Table 1. Sites Included in the Analysis, Their Geographical and Physical Characteristics, Mean (\pm SD) Loss on Ignition (LOI) and Bulk Density (BD), Sample Size (N), and Data Sources

Site	Community Type	Latitude ($^{\circ}$ N)	Tidal Amp (cm)	Loss on Ignition (g g^{-1})	Bulk Density (g cm^{-3})	N
Apcola, FL ^a	Tidal brackish marsh	29.8	25	0.43 \pm 0.12	0.14 \pm 0.07	297
FCEman, FL ^b	Everglades brackish mangrove	25.2	10	0.26 \pm 0.14	0.37 \pm 0.14	72
FCEmmm, FL ^b	Everglades mixed mangrove-marsh	25.2	5	0.17 \pm 0.15	0.35 \pm 0.13	41
FCEbrf, FL ^b	Everglades brackish tidal forested	25.4	20	0.81 \pm 0.13	0.14 \pm 0.03	39
FCEbtma, FL ^b	Everglades brackish tidal mangrove	25.4	30	0.37 \pm 0.18	0.32 \pm 0.14	80
CRMS, LA ^h	Tidal salt to tidal fresh marsh	29		0.38 \pm 0.24	0.28 \pm 0.22	500
GrandB, MS ^a	<i>S. alterniflora</i> /Juncus tidal marsh	30	22	0.12 \pm 0.04	0.39 \pm 0.15	59
WeeksB, AL ^a	Tidal brackish marsh	30.4	20	0.17 \pm 0.04	0.35 \pm 0.08	60
NorthI, SC ^a	Tidal salt marsh, <i>S. alterniflora</i>	33.3	70	0.13 \pm 0.08	0.78 \pm 0.34	278
Waccam, SC ^a	Tidal fresh marsh	33.5	60	0.55 \pm 0.16	0.13 \pm 0.09	337
Smyrna, NC ^a	Tidal brackish marsh <i>Juncus roemerianus</i>	34.8	19	0.27 \pm 0.15	0.43 \pm 0.37	416
SandH, NC ^a	Tidal brackish marsh <i>Juncus roemerianus</i>	35	5	0.56 \pm 0.20	0.23 \pm 0.26	145
James, VA ^a	Tidal fresh marsh, <i>Polygonum</i> , <i>Peltandra</i> , <i>Zizania</i> , others, & forest	37.3	40	0.21 \pm 0.06	0.44 \pm 0.13	25
Pamunk, VA ^{a,e}	Tidal fresh marsh	37.6	44	0.21 \pm 0.08	0.35 \pm 0.15	401
BrownsIsl, CA ^l	Tidal brackish marsh, <i>Schoenoplectus</i> spp.	38.0	46	0.35 \pm 0.17	0.29 \pm 0.15	113
ChinaCp, CA ^l	Tidal salt marsh, <i>Spartina foliosa</i> , <i>Sarcocornia pacifica</i> .	38.0	66	0.12 \pm 0.04	0.50 \pm 0.10	134
CoonIs, CA ^l	Tidal salt marsh, <i>S. foliosa</i> , <i>S. pacifica</i>	38.2	72	0.16 \pm 0.09	0.44 \pm 0.22	172
MonieBay, MD ^{c,d}	Tidal brackish marsh, <i>S. alterniflora</i>	38.2	15	0.27 \pm 0.18	0.49 \pm 0.49	189
PetalumaR, CA ^l	Tidal salt marsh, <i>S. foliosa</i> , <i>S. pacifica</i>	38.2	75	0.18 \pm 0.08	0.4 \pm 0.21	111

Table 1. (continued)

Site	Community Type	Latitude (°N)	Tidal Amp (cm)	Loss on Ignition (g g ⁻¹)	Bulk Density (g cm ⁻³)	N
RushRa, CA ^l	Tidal brackish marsh, <i>Schoenoplectus</i> spp.	38.2	86	0.27 ± 0.15	0.38 ± 0.19	138
Choptk, MD ^c	Tidal fresh marsh	38.7	16	0.27 ± 0.05	0.30 ± 0.07	45
PatuxR, MD ^{c,g}	Tidal fresh marsh, mixed	38.7	28	0.34 ± 0.17	0.29 ± 0.14	230
GCREW, MD ^a	Brackish tidal marsh, <i>Schoenoplectus americanus</i> , <i>Spartina patens</i>	38.9	14	0.79 ± 0.05	0.12 ± 0.02	117
OttCr, MD ^c	Tidal fresh marsh	39.4	19	0.12 ± 0.02	0.68 ± 0.19	45
DelRiv, DE ^f	Tidal fresh marsh	39.8	91	0.14 ± 0.04	0.27 ± 0.17	11
JamBay, NY ^{i,k}	<i>S. alterniflora</i> salt marsh	40.6	80	0.41 ± 0.18	0.31 ± 0.32	133
NarrBay, RI ^{aj}	Tidal salt marsh, <i>S. patens</i> , <i>D. spicata</i> , <i>S. alterniflora</i>	41.6	65	0.28 ± 0.19	0.41 ± 0.31	127
Hudson, NY ^c	Tidal fresh marsh	42	61	0.16 ± 0.1	0.58 ± 0.21	103
PlumIs, MA ^a	<i>S. patens</i> , <i>S. alterniflora</i> salt marsh	43	150	0.29 ± 0.14	0.26 ± 0.12	320
Moody, ME ^a	Tidal salt marsh, <i>S. patens</i> , <i>D. spicata</i> , <i>S. alterniflora</i>	43.2	134	0.31 ± 0.14	0.34 ± 0.22	52
Webhannet, ME ^a	Tidal salt marsh, <i>S. patens</i> , <i>D. spicata</i> , <i>S. alterniflora</i>	43.3	134	0.25 ± 0.13	0.44 ± 0.20	36
Morse, ME ^a	Tidal salt marsh, <i>S. patens</i> , <i>D. spicata</i> , <i>S. alterniflora</i>	43.7	139	0.22 ± 0.12	0.42 ± 0.21	37
Sprague, ME ^a	Tidal salt marsh, <i>S. patens</i> , <i>D. spicata</i> , <i>S. alterniflora</i>	43.7	139	0.20 ± 0.12	0.43 ± 0.24	212
Grand Means				0.30 ± 0.17	0.36 ± 0.14	5075

^aoriginal data; ^bChambers and Pederson, 2006 and <http://fcelter.fiu.edu/data/>; ^cMerrill, 1999; ^dWard et al., 1988; ^eNeubauer, 2000; ^fChurch et al., 2006; ^gOrson et al., 1992; ^hKhan and Brush, 1994; ⁱ<http://cims.coastal.la.gov/monitoring-data/>; ^jWigand et al., 2014; ^kSwanson and Wilson, 2008; ^lCallaway et al., 2012

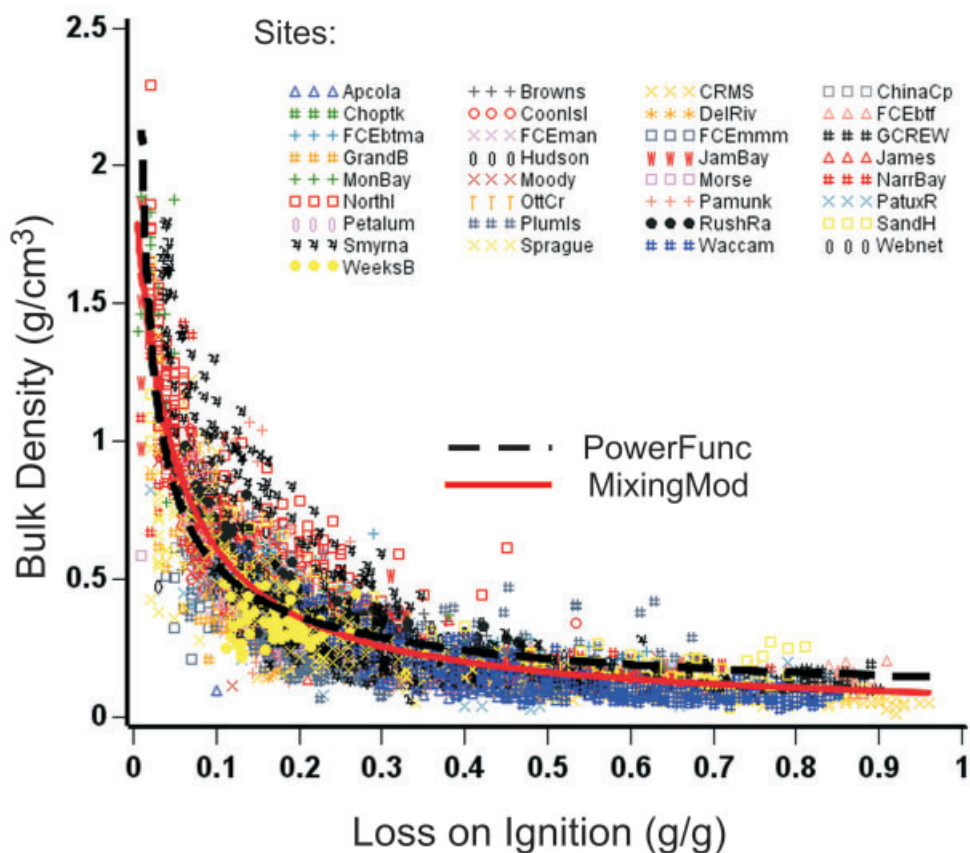


Figure 1. Best fits of the power function (— —) $BD = a LOI^b$ and ideal mixing model (—) $BD = 1/[LOI/k_1 + (1 - LOI)/k_2]$ to the complete collection of 5075 sediment samples from 33 tidal marshes and mangroves distributed around the United States. The data in Figure 1 are available in Appendix S1, Supporting Information

3.4. Model 3 Ideal Mixing Model

Stewart *et al.* [1970] suggested that in natural soils containing organic matter, there is a relationship between soil apparent density and organic matter content, namely that the bulk volume of a soil approximates to the summed self-packing volumes of the organic and mineral components. In other words, it is assumed that the bulk densities of pure organic matter and pure mineral matter are assumed constants, and in a mixture, the volumes occupied by the organic and mineral components are additive [Federer *et al.*, 1993]. For example, starting with dry weights W_o and W_i of organic and inorganic matter having bulk densities of k_1 and k_2 , and volumes $V_o = W_o/k_1$ and $V_i = W_i/k_2$, when the two are mixed, the resulting bulk density $BD = (W_o + W_i)/(W_o/k_1 + W_i/k_2)$. The mixture will have a loss on ignition $LOI = W_o/(W_o + W_i)$. Rearranging $W_o + W_i = W_o/LOI$ and substituting gives $BD = (W_o/LOI)/(W_o/k_1 + W_i/k_2)$. Substituting for W_i gives $BD = (W_o/LOI)/[W_o/k_1 + (W_o/LOI - W_o)/k_2]$, which reduces to

$$BD = 1 / [LOI/k_1 + (1 - LOI)/k_2] \tag{3}$$

This is known as the ideal mixing model [Adams, 1973]. Coefficients k_1 and k_2 are the bulk, self-packing densities of pure organic and mineral matter, respectively.

The best fit over all sites ($R^2 = 0.78$) gave k_1 and k_2 coefficients of $0.085 \pm 0.0007 \text{ g cm}^{-3}$ ($p < 0.0001$) and $1.99 \pm 0.028 \text{ g cm}^{-3}$ ($p < 0.0001$), respectively (Figure 1). Note that when $LOI = 0$, Eq. 3 reduces to $BD = k_2$, and Eq. 3 reduces to $BD = k_1$ when $LOI = 1$. When fitted specifically to each site, the resulting k_1 values varied from 0.05 to 0.168 g cm^{-3} and k_2 from 0.59 to 2.87 g cm^{-3} . Parameter estimates made from site-specific data were unreliable because few sites possessed the full spectrum of BD and LOI values needed for a good fit.

3.5. Vertical Accretion

There have been several analyses of the relative contribution of organic and inorganic accretion to vertical accretion using a top-down approach [Nyman *et al.*, 1993, 2006; Turner *et al.*, 2000; Neubauer, 2008]. Starting with known accretion rates and composition, Turner *et al.* [2000] and Neubauer [2008] calculated the densities of organic and mineral matter that best fit the total accretion rates by linear regression. From their published inventories of mineral and organic matter above the 1963 ^{137}Cs maxima, we calculated vertical accretion by dividing the mass accretion rates ($\text{g cm}^{-2} \text{ yr}^{-1}$) by the k_1 and k_2 densities, 0.085 and 1.99 g cm^{-3} , respectively, from the present study. Our calculated vertical accretion rates and their measured rates (Figure 2) were highly correlated ($r^2 = 0.6$, $p < 0.0001$).

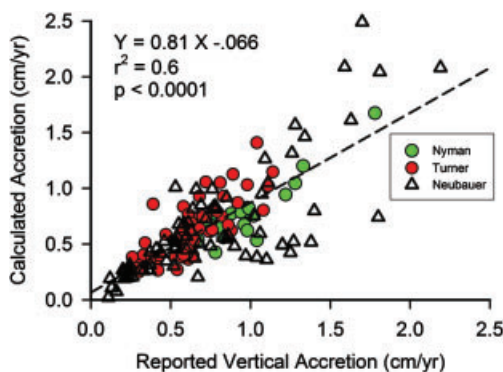


Figure 2. Mixing model-calculated accretion rates and reported ^{137}Cs -based vertical accretion rates. Calculations used published mass accretion rates ($\text{g m}^{-2} \text{ yr}^{-1}$) of inorganic and organic matter as reported in Nyman *et al.* [1993, 2006], Turner *et al.* [2000], and Neubauer [2008].

when flooded semi-diurnally (704 floods per year) by 20 cm of water at high tide would receive an annual deposit of $0.42 \text{ g cm}^{-2} \text{ yr}^{-1}$ ($=3 \times 10^{-5} \text{ g cm}^{-3} \times 20 \text{ cm} \times 704 \text{ yr}^{-1}$), using a flow-weighted, average suspended sediment concentration of 30 mg/liter for East and eastern Gulf Coast estuaries [Weston, 2014]. Dividing by k_2 (1.99 g cm^{-3}) gives 0.2 cm/yr.

As the production of the vegetation is also sensitive to flood depth [Morris *et al.*, 2002, 2013], we assumed belowground productivities of 800, 1600, and $2500 \text{ g dry weight m}^{-2} \text{ yr}^{-1}$ in high, mid, and low marsh, respectively, with 10% preservation. The organic contribution to vertical accretion is determined by the input of the refractory organic production, which should be nearly equal to the production of lignin. Using a lignin concentration of 10% [Hodson *et al.*, 1984; Wilson *et al.*, 1986; Buth and Voesenek, 1987], the vertical accretion due to a theoretical maximum belowground dry weight production of $2500 \text{ g m}^{-2} \text{ yr}^{-1}$ for *Spartina alterniflora* [Morris *et al.*, 1984] should be about $0.3 \text{ cm yr}^{-1} \approx (0.25 \text{ g cm}^{-2} \text{ yr}^{-1} \times 0.1 \text{ g lignin/g dry weight}) / 0.085 \text{ g cm}^{-3}$. Shown here (Figure 3) are the vertical accretion rate and LOI resulting from various combinations of inputs. The results show that the highest vertical accretion rates are possible only at the highest TSS concentrations, that is, the highest inputs of mineral matter (Figure 3). However, for conditions typical of East Coast estuaries (high production, low TSS), organic matter probably contributes about 3/5 or 60% to vertical accretion, which is in line with Neubauer's [2008] estimate of 62%.

4. Discussion

The variability in BD not explained by the mixing model can be due to sampling and analytical errors, as discussed earlier, and to violations of the model assumptions, namely that k_1 and k_2 are constant and their volumes additive. Additivity is not strictly true because organic molecules encapsulated within the lattices of clay minerals or sorbed to the surfaces of clay minerals [Mortland, 1970; van Veen and Kuikman, 1990; Mayer, 1994; Hedges and Keil, 1995; Bergamaschi *et al.*, 1997; Mayer *et al.*, 2004] may raise the LOI without adding to the total volume. Furthermore, a portion of the variability also could be due to compaction or differences in sediment grain size. The mass per unit volume of the solid inorganic component, that is, excluding voids and water, is commonly taken to be 2.65 g cm^{-3} [Skopp, 2000; Blanco-Canqui *et al.*, 2004].

The difference between the solid inorganic density and k_2 equates to an average porosity of 25%, but porosity and bulk density are functions of grain size [Manrique and Jones, 1991; Jackson *et al.*, 2014], which is variable. However, the fidelity of the mixing model (Figure 1) suggests that these deviations from the model assumptions are minor, at least within the universe of data analyzed here.

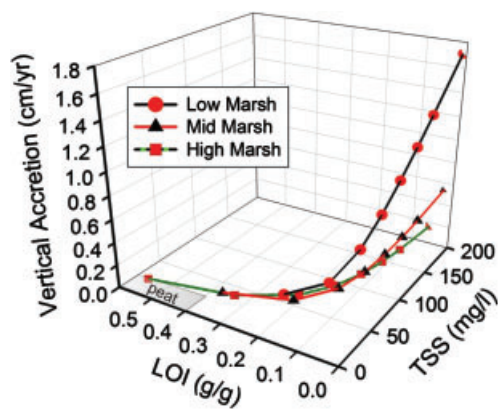


Figure 3. A curve family describing relationships among vertical accretion, soil organic content (LOI), and total suspended inorganic sediment (TSS) in hypothetical high-, mid-, and low-marsh habitats. Calculations of vertical accretion due to organic production and mineral sedimentation are based on the pure packing densities k_1 and k_2 (see text) derived from a fit of the mixing model to the entire dataset. High, mid, and low marshes were defined as having high tides that flood the surface semi-diurnally, with average water depths of 1, 5, and 20 cm, and with belowground refractory organic matter inputs of 80, 160, and 250 dry g m⁻² yr⁻¹, respectively. The grey area is the domain of peat formation, truncated at LOI = 0.6.

Mixing model parameters k_1 and k_2 were used to parse total vertical accretion into its organic and mineral fractions and to place upper limits on vertical accretion. With respect to the contributions from organic production, it can be assumed that only belowground biomass contributes to soil organic matter in tidal salt marshes because we do not observe the accumulation of a litter layer on the marsh surface. Furthermore, excluding sites of marsh restoration or newly emerging marshes where the live biomass may build over a period of years, the input of labile organic matter does not add new volume because its production and decay rate are in equilibrium, that is, the volume of living and labile organic matter exists in steady state. The refractory fraction of belowground production, however, does contribute to the change in soil volume, and we can equate this to the production of lignin, which is considered to be the most refractory part of plant tissue and is a significant component of soil organic matter in anaerobic environments such as peatlands [Morita, 1962]. Its decomposition depends primarily on the supply of oxygen [Andriess, 1988; Williams and Yavitt, 2003]. An upper limit for belowground dry weight production of 2500 g m⁻² yr⁻¹ containing 10% lignin gives a theoretical steady-state vertical accretion of about 0.3 cm yr⁻¹ (3 mm yr⁻¹). Thus, peat marshes are not likely to indefinitely survive an accelerating SLR without a source of mineral sediment.

Empirical measures of belowground production sometimes exceed the generalized estimate above 2500 g m⁻² yr⁻¹, for example, 3500 g m⁻² yr⁻¹ in Great Sippewissett marsh [Valiela *et al.*, 1976], 6500 g m⁻² yr⁻¹ in Delaware Bay [Roman and Daiber, 1984], and 7900 g m⁻² yr⁻¹ in Narragansett Bay [Wigand, 2008]. Measurements of belowground production are subject to significant sampling error [Singh *et al.*, 1984], but theoretical estimates associated with model assumptions also have uncertainty. Moreover, maximum productivity varies greatly among plant communities and species. The productivity of tidal freshwater

There was a weak, but insignificant ($p = 0.25$) decline of mean LOI by site with increasing latitude (-0.01 per degree). If the concentration of organic matter in sediment was reduced by rising temperature, perhaps due to an increase in decay rate [Kirschbaum, 1995], then we would have expected an increase in LOI with latitude. Alternatively, a decline with latitude would be consistent with a trend of decreasing above-ground productivity in salt marshes with increasing latitude [Kirwan *et al.*, 2009]. This decline with latitude is also consistent with the interpretation of Philben *et al.* [2014, 2015] who found that peat formed at more southerly latitudes and, during the Holocene Climatic Optimum, did not appear to be any more decomposed than peat formed in cooler latitudes or during subsequent cooler periods in younger sections of the cores. They suggested that decomposition in peatlands was regulated primarily by oxygen-exposure time, not by temperature.

We found a weakly negative (-0.002 cm⁻¹) but significant ($p = 0.05$) effect of tidal amplitude on LOI. Tidal amplitude could affect LOI in a number of ways, for example, by affecting dewatering and oxygen exposure, hydroperiod, and/or the effect of tidal energy on mineral sediment transport. For a fixed input of organic matter, the resulting LOI will be determined by the input of mineral matter. The latter would be controlled by the concentration of suspended sediment as well as the hydroperiod or settling time for suspended sediment.

For a fixed input of organic matter, the resulting LOI will be determined by the input of mineral matter. The latter would be controlled by the concentration of suspended sediment as well as the hydroperiod or settling time for suspended sediment.

Empirical measures of belowground production sometimes exceed the generalized estimate above 2500 g m⁻² yr⁻¹, for example, 3500 g m⁻² yr⁻¹ in Great Sippewissett marsh [Valiela *et al.*, 1976], 6500 g m⁻² yr⁻¹ in Delaware Bay [Roman and Daiber, 1984], and 7900 g m⁻² yr⁻¹ in Narragansett Bay [Wigand, 2008]. Measurements of belowground production are subject to significant sampling error [Singh *et al.*, 1984], but theoretical estimates associated with model assumptions also have uncertainty. Moreover, maximum productivity varies greatly among plant communities and species. The productivity of tidal freshwater

marshes, for example, exceeds that of salt marshes [Odum, 1988]. Primary production can also vary strongly with factors such as interannual variation in sea level [Morris *et al.*, 2002], nitrogen loading, and atmospheric CO₂ concentration [Langley *et al.*, 2009]. Therefore, belowground production might provide greater accretion than the 0.3 cm yr⁻¹ estimated here and also potentially greater rates of carbon sequestration.

The relative importance of inorganic and organic contributions to vertical accretion is dynamic over long timescales [Kearney *et al.*, 1994; Drexler, 2011] and can vary spatially in contemporary settings [Craft *et al.*, 1993]. Vertical accretion is ultimately limited by the mass inputs of mineral and refractory organic material, constrained by their bulk densities. Mass accretion rates of mineral and organic matter in tidal freshwater and salt water wetlands and ¹³⁷Cs-based vertical accretion have been reported by Nyman *et al.* [1993, 2006], Turner *et al.* [2000], and Neubauer [2008] among others, and these authors have advanced the argument that organic matter accretion is the principal means by which marshes accrete vertically. Our analysis supports this view, considering the low TSS concentrations typical of many estuaries, but we would argue that vertical accretion is limited by low availability of mineral sediment. Using a top-down approach, Turner *et al.* [2000] regressed the vertical accretion rate against organic matter accumulation and found a slope of 11 cm³ g⁻¹. Its reciprocal is k_1 and, ignoring a regression intercept, is equivalent to 0.09 g cm⁻³. His regression of vertical accretion rate against inorganic accumulation rate gave a slope of 1.36 cm³ g⁻¹ or 0.73 g cm⁻³, which is significantly less than k_2 derived in our study. Using multiple linear regression of vertical accretion rate against organic plus mineral accumulation, Neubauer [2008] derived densities equivalent to 0.182 g cm⁻³ and 0.847 g cm⁻³ for the organic and mineral fractions, respectively. The differences between coefficients derived by Turner *et al.* [2000] and Neubauer [2008] and our calculated values of k_1 and k_2 reflect the limited range of soil properties in those other studies, notably a near absence of high LOI samples.

Several generalities emerge from our bottom-up calculations of vertical accretion (Figure 3). Firstly, the theoretical maximum vertical accretion rate at steady state for a typical marsh on the East Coast and eastern Gulf is probably not greater than 0.5 cm yr⁻¹. This is the expected vertical accretion from the combined inputs of organics and inorganics at the highest loading rates, that is, complete capture of 30 mg L⁻¹ suspended sediment under 20 cm of water flooding the surface semi-diurnally and with the addition of 250 g dry weight m⁻² yr⁻¹ of refractory organic matter. However, a significantly higher sediment load, like that available in a deltaic environment, for example, 200 mg L⁻¹, could support a vertical accretion rate of 1.7 cm yr⁻¹ (Figure 3). Secondly, the lowest concentration of sediment organic matter or LOI occurs at the highest vertical accretion rates for a marsh at steady-state equilibrium with sea level, reflecting low marsh conditions with high mineral inputs. Conversely, the highest LOI occurs at the lowest accretion rates and in high marsh habitat with little flooding (Figure 3). High marsh habitat has the lowest opportunity for mineral loading from tidal flooding and, thus, little supplement to organic accretion by inorganic material.

Our analysis addresses marshes operating in steady state with typical rates of plant production and sediment availability. However, marshes often accrete sediment at different rates (Figure 2) because they are not always constrained by the limits we have assumed on mineral and organic matter accretion. Locally, there may be erosion of marsh edge and redeposition onto the remaining marsh platform [Wilson *et al.*, 2014] or episodes of rapid sedimentation following hurricanes [Cahoon *et al.*, 1995; Turner *et al.*, 2006; McKee and Cherry, 2009; Baustian and Mendelssohn, 2015]. However, storm impacts also can be negative [Chabreck and Palmisano, 1973; Cahoon, 2006; Howes *et al.*, 2010; Morton and Barras, 2011; Leonardi *et al.*, 2016], and it seems unlikely that storms could contribute globally to a positive, steady-state accretion or that they could accommodate an acceleration in sea-level rise. There are also situations when a system may be far removed from the local equilibrium state. What follows is a rapid recovery and accretion toward the equilibrium [French and Burningham, 2003]. This could happen following a wetland restoration, an emergence of a bar at the mouth of a delta, or a sudden loss in relative elevation due to storm-induced erosion or to mechanical means. These are transient conditions that could affect short-term accretion rates but that should not affect the steady-state accretion rate. Also possible are changes in state following a change in nutrient loading, which could affect organic matter production and preservation or a change in sediment loading. These could increase or decrease the steady-state accretion rate, therefore affecting the ability of tidal wetlands to maintain elevation relative to a rising sea level.

A tidal wetland with a high LOI and low accretion rate is not necessarily at its maximum rate of vertical accretion because the mineral loading rate and primary productivity will depend on its relative

elevation, depth of flooding, and suspended sediment concentration. With an accelerated rate of SLR, the relative, equilibrium marsh elevation will fall, and the vertical accretion rate will rise [Morris *et al.*, 2002]. It is only when the equilibrium elevation falls below the optimum for primary production that the marsh will fail to keep up with the acceleration in SLR. Our analysis, which assumes only chronic inputs, suggests that this will most likely occur when the local rate of SLR (including subsidence) exceeds about 0.5 cm yr^{-1} in most East Coast estuaries or 1.7 cm yr^{-1} in a sediment-rich estuary. Thus, even current rates of SLR along the East Coast (mean = 0.34 cm yr^{-1} , range = $0.18\text{--}0.60 \text{ cm yr}^{-1}$) (<http://tidesandcurrents.noaa.gov/sltrends/sltrends.html>) are near the point that will lead to long-term elevation loss and eventual drowning of coastal wetlands.

Acknowledgments

This research was supported by NOAA grant NA10NOS4780146, NSF DEB grants 1052636, 0620409, 9910514, 1237517, 1127958, 0950090, 1457100 and 1355059, NSF EAR grants 1322859 and 1322658, NSF OCE-1238212, NICCR grant DE-FC02-06ER64298, NASA grant NNN14AY67, the Bates College Student Research Fund, the Bryn Mawr College Summer Science Research Program, and the Blue Carbon Initiative coordinated by Conservation International, IUCN, and IOC-UNESCO.

References

- Adams, W. A. (1973), The effect of organic matter on the bulk and true densities of some uncultivated podzolic soils, *J. Soil Sci.*, *24*, 10–17.
- Andriessse, J. P. (1988), Nature and management of tropical peat soils, *FAO Soils Bull.* *59*, Food and Agriculture Organization United Nations, Rome.
- Baustian, J. J., and I. A. Mendelssohn (2015), Hurricane-induced sedimentation improves marsh resilience and vegetation vigor under high rates of relative sea level rise, *Wetlands*, *35*, 795–802.
- Bergamaschi, B. A., E. Tsamakidis, R. G. Keil, T. I. Eglinton, D. B. Montluçon, and J. I. Hedges (1997), The effect of grain size and surface area on organic matter, lignin and carbohydrate concentration, and molecular compositions in Peru Margin sediments, *Geochim. Cosmochim. Acta*, *61*, 1247–1260.
- Blanco-Canqui, H., C. J. Gantzer, S. H. Anderson, and E. E. Alberts (2004), Tillage and crop influences on physical properties for an Epiaqualf, *Soil Sci. Soc. Am. J.*, *68*, 567–576.
- Buth, G. J. C., and L. A. C. J. Voesenek (1987), Decomposition of standing and fallen litter of halophytes in a Dutch salt marsh, in *Geobotany 11: Vegetation Between Land and Sea*, edited by A. H. L. Huiskes, C. W. P. M. Blom, and J. Rozema, pp. 146–165, Dr. W. Junk Pub, Dordrecht, The Netherlands.
- Cahoon, D. R., D. J. Reed, J. W. Day Jr., G. D. Steyer, R. M. Boumans, J. C. Lynch, D. McNally, and N. Latif (1995), The influence of Hurricane Andrew on sediment distribution in Louisiana coastal marshes, *J. Coastal Res.*, *21*, 280–294.
- Cahoon, D. R. (2006), A review of major storm impacts on coastal wetland elevations, *Estuaries Coasts*, *29*, 889–898.
- Callaway, J. C., E. L. Borgnis, R. E. Turner, and C. S. Milan (2012), Carbon sequestration and sediment accretion in San Francisco Bay tidal wetlands, *Estuar. Coast.*, *35*, 1163–1181.
- Chabreck, R. H., and A. W. Palmisano (1973), The effects of Hurricane Camille on the marshes of the Mississippi River Delta, *Ecology*, *54*, 1118–1123.
- Chambers, R. M., and K. A. Pederson (2006), Variation in soil phosphorus, sulfur, and iron pools among south Florida wetlands, *Hydrobiologia*, *569*, 63–70.
- Church, T. M., C. K. Sommerfield, D. J. Velinsky, D. Point, C. Benoit, D. Amouroux, D. Plaa, and O. F. X. Donard (2006), Marsh sediments as records of sedimentation, eutrophication and urban pollution in the urban Delaware Estuary, *Mar. Chem.*, *102*, 72–95.
- Craft, C., J. Clough, J. Ehman, S. Joye, R. Park, S. Pennings, H. Guo, and M. Machmuller (2009), Forecasting the effects of accelerated sea-level rise on tidal marsh ecosystem services, *Frontiers Ecol.*, *7*, 73–78.
- Craft, C. B., E. D. Senica, and S. W. Broome (1993), Vertical accretion in microtidal regularly and irregularly flooded estuarine marshes, *Estuarine Coastal Shelf Sci.*, *37*, 371–386.
- Davis, J. L., C. A. Currin, C. O'Brien, C. Raffenburg, and A. Davis (2015), Living shorelines: coastal resilience with a blue carbon benefit, *PLoS One*, *10*(11), e0142595, doi:10.1371/journal.pone.0142595.
- Dean, W. E., Jr. (1974), Determination of carbonate and organic matter in calcareous sediments and sedimentary rocks by loss on ignition: Comparison with other methods, *J. Sed. Res.*, *44*, 242–248.
- Deegan, L. A., D. S. Johnson, R. S. Warren, B. J. Peterson, J. W. Fleeger, S. Fagherazzi, and W. M. Wollheim (2012), Coastal eutrophication as a driver of saltmarsh loss, *Nature*, *490*, 388–392.
- Drexler, J. Z. (2011), Peat formation processes through the millennia in tidal marshes of the Sacramento-San Joaquin Delta, California, USA, *Estuaries Coasts*, *34*, 900–911.
- Engelhart, S. E., and B. P. Horton (2012), Holocene sea level database for the Atlantic coast of the United States, *Quat. Sci. Rev.*, *54*, 12–25.
- Engelhart, S. E., M. Vacchi, B. P. Horton, A. R. Nelson, and R. E. Kopp (2015), A sea-level database for the Pacific coast of central North America, *Quat. Sci. Rev.*, *113*, 78–92, doi:10.1016/j.quascirev.2014.12.001.
- Federer, C. A., D. E. Turcotte, and C. T. Smith (1993), The organic fraction–bulk density relationship and the expression of nutrient content in forest soils, *Can. J. Forest Res.*, *23*, 1026–1032.
- French, J. R., and H. Burningham (2003), Tidal marsh sedimentation versus sea-level rise: a southeast England estuarine perspective, in *Proceedings International Conference on Coastal Sediments 2003*, edited by R. A. Davis, pp. 1–14, World Scientific Publishing and East Meets West Productions, Corpus Christi, Tex.
- Gosselink, J. G., R. Hatton, and C. S. Hopkins (1984), Relationship of organic carbon and mineral content to bulk density in Louisiana marsh soils, *Soil Sci.*, *137*, 177–180.
- Grigal, D. F., S. L. Brovold, W. S. Nord, and L. F. Ohmann (1989), Bulk density of surface soils and peat in the North Central United States, *Can. J. Soil Sci.*, *69*, 895–900.
- Harrison, A. F., and K. L. Bockock (1981), Estimation of soil bulk-density from loss-on-ignition values, *J. Appl. Ecol.*, *8*, 919–927.
- Hedges, J. I., and R. G. Keil (1995), Sedimentary organic matter preservation: an assessment and speculative synthesis, *Mar. Chem.*, *49*, 81–115.
- Heiri, O., A. F. Lotter, and G. Lemcke (2001), Loss on ignition as a method for estimating organic and carbonate content in sediments: Reproducibility and comparability of results, *J. Paleolimnol.*, *25*, 101–110.
- Hodson, R. E., R. R. Christian, and A. E. Maccubbin (1984), Lignocellulose and lignin in the salt marsh grass *Spartina alterniflora*: Initial concentrations and short-term, post-depositional changes in detrital matter, *Marine Biol.*, *81*, 1–7.
- Honeysett, J. L., and D. A. Ratkowsky (1989), The use of ignition loss to estimate bulk density of forest soils, *J. Soil Sci.*, *40*, 1365–2389.

- Horton, B. P., S. Rahmstorf, S. E. Engelhart, and A. C. Kemp (2014), Expert assessment of sea-level rise by AD 2100 and AD 2300, *Quat. Sci. Rev.*, *84*, 1–6.
- Howes, N. C., D. M. FitzGerald, Z. J. Hughes, I. Y. Georgiou, M. A. Kulp, M. D. Miner, J. M. Smith, and J. A. Barras (2010), Hurricane-induced failure of low salinity wetlands, *Proc. Natl. Acad. Sci. U. S. A.*, *210*, 14014–14019.
- Jackson, C. R., J. A. Thompson, and R. K. Kolka (2014), Wetland soils, hydrology and geomorphology, Chapter 2, in *Ecology of Freshwater and Estuarine Wetlands*, edited by D. Batzer and R. Sharitz, pp. 23–60, University of California Press, Berkeley.
- Jeffrey, D. W. (1970), A note on the use of ignition loss as a means for the approximate estimation of soil bulk density, *J. Ecol.*, *58*, 297–299.
- Kearney, M. S., J. C. Stevenson, and L. G. Ward (1994), Spatial and temporal changes in marsh vertical accretion rates at Monie Bay: Implications for sea-level rise, *J. Coastal Res.*, *10*, 1010–1020.
- Kelley, J. T., D. F. Belknap, G. L. Jacobson, and H. A. Jacobson (1988), The morphology and origin of salt marshes along the glaciated coastline of Maine, USA, *J. Coastal Res.*, *4*, 649–665.
- Kemp, A. C., B. P. Horton, J. P. Donnelly, M. E. Mann, M. Vermeer, and S. Rahmstorf (2011), Climate related sea-level variations over the past two millennia, *Proc. Natl. Acad. Sci. U. S. A.*, *108*, 11017–11022, doi:10.5670/oceanog.2011.28.
- Kemp, A. C., B. P. Horton, C. H. Vane, C. E. Bernhardt, D. R. Corbett, S. E. Engelhart, S. C. Anisfeld, A. C. Parnell, and N. Cahill (2013), Sea-level change during the last 2500 years in New Jersey, USA, *Quat. Sci. Rev.*, *81*, 90–104, doi:10.1016/j.quascirev.2013.09.024.
- Khan, H., and G. S. Brush (1994), Nutrient and metal accumulation in a freshwater tidal marsh, *Estuaries*, *17*, 345–360.
- Kirschbaum, M. U. F. (1995), The temperature dependence of soil organic matter decomposition, and the effect of global warming on soil organic storage, *Soil Biol. Biochem.*, *27*, 753–760.
- Kirwan, M. L., G. R. Guntenspergen, and J. T. Morris (2009), Latitudinal trends in *Spartina alterniflora* productivity and the response of coastal marshes to global change, *Global Change Biol.*, *15*, 1982–1989.
- Kirwan, M. L., and J. P. Megonigal (2013), Tidal wetland stability in the face of human impacts and sea-level rise, *Nature*, *504*, 53–60, doi:10.1038/nature12856.
- Kirwan, M. L., G. R. Guntenspergen, A. D'Alpaos, J. T. Morris, S. M. Mudd, and S. Temmerman (2010), Limits on the adaptability of coastal marshes to rising sea level, *J. Geophys. Res.*, *37*, L23401, doi:10.1029/2010GL045489.
- Kirwan, M. L., A. B. Murray, J. P. Donnelly, and D. R. Corbett (2011), Rapid wetland expansion during European settlement and its implication for marsh survival under modern sediment delivery rates, *Geology*, *39*, 507–510, doi:10.1130/G31789.1.
- Krone, R. B. (1987), A method for simulating historic marsh elevations, in *Coastal Sediments '87*, edited by N. C. Krause, pp. 316–323, New York, American Society of Civil Engineers.
- Langley, J. A., K. L. McKee, D. R. Cahoon, J. A. Cherry, and J. P. Megonigal (2009), Elevated CO₂ stimulates marsh elevation gain, counterbalancing sea-level rise, *Proc. Natl. Acad. Sci. U. S. A.*, *106*, 6182–6186, doi:10.1073/pnas.0807695106.
- Leonardi, N., N. K. Ganju, and S. Fagherazzi (2016), A linear relationship between wave power and erosion determines salt-marsh resilience to violent storms and hurricanes, *Proc. Natl. Acad. Sci. U. S. A.*, *113*, 64–68, doi:10.1073/pnas.1510095112.
- Levenberg, K. (1944), A method for the solution of certain non-linear problems in least squares, *Q. Appl. Math.*, *2*, 164–168.
- Manrique, L. A., and C. A. Jones (1991), Bulk density of soils in relation to soil physical and chemical properties, *Soil Sci. Soc. Am. J.*, *55*, 476–481.
- Mariotti, G., and S. Fagherazzi (2013), Critical width of tidal flats triggers marsh collapse in the absence of sea-level rise, *Proc. Natl. Acad. Sci. U. S. A.*, *110*, 5353–5356, doi:10.1073/pnas.1219600110.
- Mayer, L. M. (1994), Relationships between mineral surfaces and organic carbon concentrations in soils and sediments, *Chem. Geol.*, *114*, 347–363.
- Mayer, L. M., L. L. Schick, K. R. Hardy, R. Wagai, and J. McCarthy (2004), Organic matter in small mesopores in sediments and soils, *Geochim. Cosmochim. Acta*, *68*, 3863–3872.
- McKee, K. L., and J. A. Cherry (2009), Hurricane Katrina sediment slowed elevation loss in subsiding brackish marshes of the Mississippi River Delta, *Wetlands*, *29*, 2–15.
- Merrill, J. Z. (1999), Tidal freshwater marshes as nutrient sinks: Particulate nutrient burial and denitrification, PhD dissertation, Univ. of Maryland.
- Morita, H. (1962), Composition of peat humus and its derivatives. Oxidative conversion to lignin model compounds, *J. Org. Chem.*, *27*, 1079–1080.
- Morris, J. T., and W. B. Bowden (1986), A mechanistic, numerical model of sedimentation, mineralization, and decomposition for marsh sediments, *Soil. Sci. Soc. Am. J.*, *50*, 96–105.
- Morris, J. T., R. A. Houghton, and D. B. Botkin (1984), Theoretical limits of belowground production by *Spartina alterniflora*: An analysis through modelling, *Ecol. Modell.*, *26*, 155–175.
- Morris, J. T., P. V. Sundareswar, C. T. Nietch, B. Kjerfve, and D. R. Cahoon (2002), Responses of coastal wetlands to rising sea level, *Ecology*, *83*, 2869–2877.
- Morris, J. T., K. Sundberg, and C. S. Hopkinson (2013), Salt marsh primary production and its responses to relative sea level and nutrients in estuaries at Plum Island, Massachusetts, and North Inlet, South Carolina, USA, *Oceanography*, *26*, 78–84.
- Mortland, M. M. (1970), Clay-organic complexes and interactions, *Adv. Agron.*, *22*, 75–117.
- Morton, R. A., and J. A. Barras (2011), Hurricane impacts on coastal wetlands: A half-century record of storm-generated features from Southern Louisiana, *J. Coastal Res.*, *27*, 27–43.
- Neubauer, S. C. (2000), Carbon dynamics in a tidal freshwater marsh, PhD dissertation, College of William and Mary, Williamsburg, Va.
- Neubauer, S. C. (2008), Contributions of mineral and organic components to tidal freshwater marsh accretion, *Estuarine Coastal Shelf Sci.*, *78*, 78–88.
- Nyman, J. A., R. D. Delaune, R. J. Walters, and W. H. Patrick Jr. (1993), Relationship between vegetation and soil formation in a rapidly submerging coastal marsh, *Mar. Ecol. Prog. Ser.*, *96*, 269–279.
- Nyman, J. A., R. J. Walters, R. D. Delaune, and W. H. Patrick Jr. (2006), Marsh vertical accretion via vegetative growth, *Estuarine Coastal Shelf Sci.*, *69*, 370–380.
- Odum, W. E. (1988), Comparative ecology of tidal freshwater and salt marshes, *Annu. Rev. Ecol. Syst.*, *19*, 147–176.
- Orson, R. A., R. L. Simpson, and R. E. Good (1992), A mechanism for the accumulation and retention of heavy metals in tidal freshwater marshes of the upper Delaware River estuary, *Estuarine Coastal Shelf Sci.*, *34*, 171–186.
- Percival, J. B., and P. J. Lindsay (1997), Measurement of physical properties of sediments (Chapter 2), in *Manual of Physico-Chemical Analysis of Aquatic Sediments*, edited by A. Mudroch, J. M. Azcue, and P. Mudroch, pp. 7–45, CRC Lewis Publishers, Boca Raton, Fla.
- Périé, C., and R. Ouime (2008), Organic carbon, organic matter and bulk density relationships in boreal forest soils, *Can. J. Soil Sci.*, *88*, 315–325.

- Philben, M., K. Kaiser, and R. Benner (2014), Does oxygen exposure time control the extent of organic matter decomposition in peatlands? *J. Geophys. Res. Biogeosci.*, *119*, 897–909, doi:10.1002/2013JG002573.
- Philben, M., J. Holmquist, G. MacDonald, D. Duan, K. Kaiser, and R. Benner (2015), Temperature, oxygen, and vegetation controls on decomposition in a James Bay peatland, *Global Biogeochem. Cycles*, *29*, 729–743, doi:10.1002/2014GB004989.
- Ratliff, K. M., A. E. Braswell, and M. Marani (2015), Spatial response of coastal marshes to increased atmospheric CO₂, *Proc. Natl. Acad. Sci. U. S. A.*, *112*, 15580–15584, doi:10.1073/pnas.1516286112.
- Redfield, A. C., and M. Rubin (1962), The age of salt marsh peat and its relation to recent changes in sea level at Barnstable Massachusetts, *Proc. Natl. Acad. Sci. U. S. A.*, *48*, 1728–1734.
- Roman, C. T., and F. C. Daiber (1984), Aboveground and belowground primary production dynamics of two Delaware Bay tidal marshes, *Bull. Torrey Bot. Club*, *3*, 34–41.
- Singh, J. S., W. K. Lauenroth, H. W. Hunt, and D. M. Swift (1984), Bias and random errors in estimators of net root production: A simulation approach, *Ecology*, *65*, 1760–1764.
- Skopp, J. M. (2000), Physical properties of primary particles, in *Handbook of Soil Science*, edited by M. E. Sumner, pp. A1–A17, CRC Press, Boca Raton, Fla.
- Stewart, V. I., W. A. Adams, and H. H. Abdulla (1970), Quantitative pedological studies on soils derived from Silurian mudstones, *J. Soil Sci.*, *21*, 248–255.
- Swanson, R. L., and R. E. Wilson (2008), Increased tidal ranges coinciding with Jamaica Bay development contribute to marsh flooding, *J. Coastal Res.*, *24*, 1565–1569.
- Turner, R. E., E. M. Swenson, and C. S. Milan (2000), Organic and inorganic contributions to vertical accretion in salt marsh sediments, in *Concepts and Controversies in Tidal Marsh Ecology*, edited by M. P. Weinstein and D. A. Kreeger, pp. 583–595, Kluwer Academic Publishers, Berlin, Germany.
- Turner, R. E., J. J. Baustian, E. M. Swenson, and J. S. Spicer (2006), Wetland sedimentation from Hurricanes Katrina and Rita, *Science*, *314*, 449–452.
- Valiela, I., J. M. Teal, and N. Y. Persson (1976), Production and dynamics of experimentally enriched salt marsh vegetation: belowground biomass, *Limnol. Oceanogr.*, *21*, 245–252.
- Van Veen, J. A., and P. J. Kuikman (1990), Soil structural aspects of decomposition of organic matter by micro-organisms, *Biogeochemistry*, *11*, 213–233.
- Vermeer, M., and S. Rahmstorf (2009), Global sea level linked to global temperature, *Proc. Natl. Acad. Sci. U. S. A.*, *106*, 21527–21532, doi:10.1073/pnas.0907765106.
- Ward, L. G., M. S. Kearney, and J. C. Stevenson (1988), Assessment of Marsh Stability at the Estuarine Sanctuary Site at Monie Bay: Implications for Management. Final Rep., National Oceanic and Atmospheric Administration, National Ocean Service, Office of Ocean Resource Management, Sanctuary Program Division, Washington, D. C.
- Weston, N. B. (2014), Declining sediments and rising seas: An unfortunate convergence for tidal wetlands, *Estuaries Coasts*, *37*, 1–23.
- Wigand, C. (2008), Coastal salt marsh community change in Narragansett Bay in response to cultural eutrophication (Chapter 17), in *Science for Ecosystem-based Management: Narragansett Bay in the 21st Century*, edited by B. A. Costa Pierce and A. Desbonnet, pp. 499–522, Springer Publishing, New York.
- Wigand, C., R. McKinney, M. Chintala, M. Charpentier, and G. Thursby (2003), Relationships of nitrogen loadings, residential development, and physical characteristics with plant structure in New England salt marshes, *Estuaries*, *26*, 1494–1504.
- Wigand, C., et al. (2014), Below the disappearing marshes of an urban estuary: Historic nitrogen trends and soil structure, *Ecol. Appl.*, *24*, 633–649.
- Williams, C. J., and J. B. Yavitt (2003), Botanical composition of peat and degree of peat decomposition in three temperate peatlands, *Ecoscience*, *10*, 85–95.
- Wilson, C. A., Z. J. Hughes, D. M. FitzGerald, C. S. Hopkinson, V. Valentine, and A. S. Kolker (2014), Saltmarsh pool and tidal creek morphodynamics: Dynamic equilibrium of northern latitude saltmarshes? *Geomorphology*, *213*, 99–115.
- Wilson, J. O., R. Buchsbaum, I. Valiela, and T. Swain (1986), Decomposition in salt marsh ecosystems: phenolic dynamics during decay of litter of *Spartina alterniflora*, *Mar. Ecol. Prog. Ser.*, *29*, 177–187.